

21 **Ozone pollution mitigation strategy informed by long-term**
22 **trends of atmospheric oxidation capacity**

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60 **Abstract:** Tropospheric ozone pollution is a critical air quality concern in China.
61 However, the most effective mitigation approach remains unclear, with prioritizing the
62 reduction of volatile organic compounds or nitrogen oxides (NO_x) currently still under
63 debate. Here, we analyze observational measurements of ozone in August, and its
64 precursors, from urban Beijing between 2006 and 2020. We show that despite a
65 continuous increase of the primary atmospheric oxidant (hydroxyl radical, OH), ozone
66 increased until 2014 and then decreased. This ozone trend can be explained by changes
67 in OH turnover rate, primarily determined by the reactivity ratio between volatile
68 organic compounds and NO_x. Overall, reactive abatement of volatile organic
69 compounds should be a near-future priority for ozone pollution control in China,
70 followed by further NO_x controls.

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82 During past decades China experienced rapid industrialization, which brought
83 economic advantages, but also caused significantly deteriorated air quality ^{1, 2}.
84 According to WHO guidelines, ozone (O₃) and fine particulate matter with a diameter
85 of less than 2.5 micrometers (PM_{2.5}) are the chief hazards of China's air pollution ³.
86 Over the past decade, China launched the toughest-ever clean air policy in the country
87 ⁴. These efforts, by reducing emissions of primary pollutants such as VOCs, NO_x, SO₂
88 and others, improved PM_{2.5} air quality ^{4, 5}. However, O₃ continued to increase and only
89 began to stabilize at a later stage ^{2, 6}. This contrasts with the generally decreasing O₃
90 trend in Europe and North America due to continuous reductions of VOCs and NO_x
91 emissions since the 1990s ^{7, 8, 9, 10}. The causes of the contrasting trends and adverse O₃
92 pollution trends in China are still not well understood, representing a grand challenge
93 for O₃ pollution control and development of effective strategies for continued air quality
94 improvement. Several studies have devoted efforts to explore the causes of rising O₃
95 levels in Beijing, a typical megacity in China. Wang and Zhang et al reported that the
96 local photochemistry driven by VOCs and NO_x emissions plays a vital role in
97 determining O₃ trend in the city ^{11, 12, 13}. Other research has evaluated the impact of
98 meteorological variability on the long-term trend of O₃ in Beijing ^{14, 15}, showing a minor
99 effect of meteorological variability compared to variations in anthropogenic emissions.

100 Tropospheric O₃ is directly produced from photolysis of nitrogen dioxide (NO₂).
101 NO₂ can be produced during the gas-phase oxidation of NO by peroxy radicals, which
102 are produced from the oxidation of VOCs by hydroxyl (OH) radicals ¹⁶. We
103 demonstrate that the dependence of OH radical cycling on ambient NO_x and VOCs

104 concentrations is key to understanding the O₃ trend in China and the contrast in O₃
105 trends between China and other regions.

106 **Trends of O₃ and its precursors in Beijing**

107 Beijing is one of the megacities in China, suffering from severe O₃ pollution. The
108 implementation of emission reductions in Beijing began earlier and has been more
109 stringent than in other cities in China¹⁷, which provides a good opportunity to study
110 the response of O₃ pollution to precursors' abatements from a long-term perspective.
111 The experience of Beijing can serve as a model for other cities in China that still suffer
112 from rapidly increasing O₃. This study is based on August measurements in Beijing
113 from 2006 to 2020, which is the most comprehensive, long-term observational data set
114 of O₃ and its precursors VOCs and NO_x that is available in China (Methods M1;
115 Extended data Table 1). Although the O₃ concentrations in August are not the highest
116 in summer, they are not much different from other months of the O₃ pollution season
117 (May to August) (Extended data Fig. 1). The highest O₃ concentration occurred in June
118 and July with mean maximum daily 8-hour average (MDA8) O₃ concentration of 80
119 ppbv during 2014-2020, which is higher than that in August by 9 ppbv.

120 Given that variations in meteorological and other physical conditions can affect
121 O₃ trends, we use a multiple linear regression method to remove the impact of
122 photolysis frequencies, temperature, relative humidity, wind speed and wind direction
123 on O₃, thereby obtaining a normalized O₃ (Methods M2)^{18, 19}. Figure 1B shows that
124 mean MDA8 O₃ concentrations in August increased before 2014, and then decreased in

125 the following years. The normalized O₃ also shows a non-monotonic trend, with a
126 smaller increase before 2014, compared to the mean MDA8 O₃ concentrations (Fig.
127 1D). Similar variable trends are also found for other secondary oxidation products such
128 as formaldehyde (Extended data Fig. 2).

129 Figure 1A shows the changes in OH reactivity of NO_x and VOCs (NO_x^R and
130 VOC^R) based on August measurements in Beijing from 2006 to 2020. NO_x^R and VOC^R
131 are defined as the concentration of NO_x and VOCs multiplied by their respective
132 reaction rate coefficients with OH (CO is included in the calculation of VOC^R as
133 detailed in Supplementary S2). Both NO_x^R and VOC^R show consistently decreasing
134 trends, confirming the effects of emission controls reported in earlier studies^{4, 20}.
135 Clearly, the trend of O₃ precursors is distinct from that of O₃.

136 To understand the relationship between OH radical cycling and the O₃ trend, we
137 first investigate the ratio VOC^R/NO_x^R where VOC^R and NO_x^R represent the main
138 production and sink terms of OH radicals, respectively^{21, 22}. VOC^R acts as a production
139 term because the reaction of VOCs with OH produces peroxy radicals and oxygenated
140 VOCs (OVOCs) which will increase OH through recycling and amplification of OH²³,
141²⁴. It is notable that the trend of VOC^R/NO_x^R (Fig. 1C) is similar to that of normalized
142 O₃; both remained approximately constant before 2014, and then decreased rapidly after
143 2014. The rapid decrease in VOC^R/NO_x^R after 2014 is attributed to decreases in
144 industrial emissions with high emission ratios of VOC to NO_x and in solvent usage¹⁷
145 under the more rigorous emissions standards implemented during the clean air action
146 plan in 2013-2020^{4, 20}. As shown in Fig. 1A, compared to an almost constant decreasing

147 trend of NO_x^R ($-3.4\% \text{ yr}^{-1}$, $r^2=0.66$), VOC^R decreased faster after 2014 ($-7.6\% \text{ yr}^{-1}$,
148 $r^2=0.68$) than before 2014 ($-3.0\% \text{ yr}^{-1}$, $r^2=0.45$). In addition to Beijing, the similarity of
149 trends between $\text{VOC}^R/\text{NO}_x^R$ and normalized O_3 is also seen in other Chinese urban
150 areas such as Shanghai in the Yangtze River Delta and Hong Kong in the Pearl River
151 Delta (Extended data Fig. 3); however, in contrast to Beijing these two variables have
152 consistently increased in Shanghai from 2011 to 2019 and in Hong Kong from 2013 to
153 2019, rather than showing recent decreases.

154 **The relationship between OH radical cycling and the O_3 trend**

155 To gain a deeper insight, we investigated the detailed radical cycling and
156 photochemistry by applying an advanced box model based on the updated Master
157 Chemical Mechanism v3.3 chemistry scheme (MCM 3.3)²⁵. This model is constrained
158 by the comprehensive observations and accounts for the diurnal evolution of the
159 planetary boundary layer²⁶ and heterogeneous uptake of HO_2 radicals and reactive
160 nitrogen species onto aerosols (Methods M3). The model accurately simulates the
161 observed O_3 (Fig. 1B), OH (Fig. S7) and other photochemical products (Fig. S6 and
162 Fig. S10) with biases smaller than 20%.

163 The box model simulation shows that OH increased continuously during 2006-
164 2020, giving an approximate doubling (Fig. 2A). Besides the box model simulations,
165 we have also adopted another approach, using a measurable proxy, the ratio of daytime
166 methyl vinyl ketone plus methacrolein (MVK+MACR) to isoprene^{27, 28} which is
167 approximately proportional to the OH concentration (Supplementary S5), to represent

168 the OH trend. As shown in Fig. 2A, the observed and simulated
169 (MVK+MACR)/isoprene agree, and display an increasing trend from 2010 to 2020,
170 providing strong support for the modeled OH trend. To identify the primary causes of
171 the continuous OH increase, we performed sensitivity analyses of the relevant factors.
172 The main drivers of the OH change are the increase of photolysis frequencies²⁹,
173 decrease in NO_x emissions, decrease in VOC emissions, and reduced heterogeneous
174 uptake³⁰, which leads to relative changes of +44%, +56%, -23% and +15%,
175 respectively (Fig. 2B). As the photochemical conditions over this period (2006 to 2020)
176 were always in the NO_x-saturated regime (Methods M4), NO_x reduction increased OH,
177 and VOC reduction decreased OH (Extended data Fig. 4). The combined effect of other
178 factors (including the changes in O₃, H₂O and temperature) played only a minor role in
179 the OH budget change.

180 Given the trend of OH radicals, we can approximate the ozone production rate
181 P(O₃) by Eq (1) and Eq (2), as detailed in Methods M4:

$$182 \quad P(\text{O}_3) = \alpha k_{\text{HO}_2+\text{NO}} [\text{HO}_2] [\text{NO}] \quad (1)$$

$$183 \quad \alpha k_{\text{HO}_2+\text{NO}} [\text{HO}_2] [\text{NO}] = \alpha k_{\text{OH}+\text{VOC}} [\text{OH}] [\text{VOC}] = \alpha [\text{OH}] \text{VOC}^{\text{R}} \quad (2)$$

184 where the coefficient α is equal to $([\text{RO}_2]+[\text{HO}_2])/[\text{HO}_2]$, which remained
185 approximately constant over the 2006-2020 period according to the model output
186 (Methods M4). Equation (2) holds under conditions when the radical chain cycling is
187 much more important than chain initiation^{22, 31}; these conditions are established in
188 Beijing (Methods M4; Extended data Fig. 5).

189 Figure 2C shows the trend of O₃ production calculated from the OH turnover rate

190 induced by VOCs ($[\text{OH}] \times \text{VOC}^{\text{R}}$) according to Eq (2). Consistent with the trend of
191 observed O_3 concentrations (Fig. 1B), $[\text{OH}] \times \text{VOC}^{\text{R}}$ also shows two distinct trends,
192 increasing before 2014 and decreasing after 2014. The strong positive correlation of
193 $[\text{OH}] \times \text{VOC}^{\text{R}}$ with O_3 and O_x ($\text{O}_x = \text{O}_3 + \text{NO}_2$) concentrations (Fig. 2D) also demonstrates
194 that the change of ambient O_3 over the entire 2006 to 2020 period is consistent with our
195 understanding of O_3 photochemistry and OH cycling. Given that O_3 production
196 primarily depends on $[\text{OH}] \times \text{VOC}^{\text{R}}$, the contrasting O_3 trends before and after 2014 can
197 be explained by a reduction in VOC^{R} that was smaller than the increase in OH before
198 2014, but larger after 2014.

199 Figures 1 and 2 show that by affecting OH cycling, $\text{VOC}^{\text{R}}/\text{NO}_x^{\text{R}}$ controls the
200 trends of $\text{P}(\text{O}_3)$ and O_3 concentration. The underlying mechanisms, however, remain
201 under debate. A key parameter here is the radical chain length (ChL), defined as
202 $\alpha k_{\text{HO}_2+\text{NO}}[\text{HO}_2][\text{NO}]/\text{P}(\text{RO}_x)$, which characterizes the average number of cycles each
203 RO_x radical makes prior to termination. Tonnesen and Dennis³² suggested that the
204 effect of VOC/NO_x on $\text{P}(\text{O}_3)$ changes is mainly through changing the ChL, while
205 several studies suggested its effects are mainly on primary production of RO_x
206 ($\text{RO}_x = \text{OH} + \text{HO}_2 + \text{RO}_2$) radicals ($\text{P}(\text{RO}_x)$)^{23, 33}. $\text{P}(\text{RO}_x)$ has contributions from the
207 photolysis of radical precursors (including carbonyls, nitrous acid and O_3) and alkene
208 ozonolysis (Methods M4). ChL primarily depends on the levels of VOCs and NO_x ³⁴.

209 To elucidate the actual mechanism in Beijing, we investigated the effects of both
210 ChL and $\text{P}(\text{RO}_x)$ as shown in Eq (3), where $\text{P}(\text{O}_3)$ is expressed as the product of $\text{P}(\text{RO}_x)$
211 and ChL:

212
$$P(O_3) = P(RO_X) \times ChL \quad (3)$$

213 As shown in Fig. 3A, the $P(RO_X)$ increased only slightly before 2014 but decreased
214 after 2014 with a trend similar to that of $[OH] \times VOC^R$ (i.e., $P(O_3)$), while the ChL level
215 remained stable with no significant change. This is clear evidence that in Beijing it is
216 the change of $P(RO_X)$ rather than ChL that controls the distinct trends of O_3 production
217 before and after 2014. Among the different contributors to $P(RO_X)$, photolysis of
218 oxygenated VOCs (OVOCs) was the largest contributor (Fig. 3B). Both measurements
219 and model simulations show a strong dependence of OH reactivity of OVOCs ($OVOC^R$)
220 on VOC^R/NO_X^R (Fig. 3C). The decreasing VOC^R/NO_X^R after 2014 (Fig. 1C) weakened
221 OH radical cycling, resulting in a decrease in the production of OVOCs, which in turn
222 reduced $P(RO_X)$ to further reduce the production of O_3 (Fig. 3D). Note that reduction
223 of VOCs also contributed to the reduction of OVOCs but is less important than the
224 reduction of VOC^R/NO_X^R . The source apportionment of OVOCs indicates that
225 reductions in primary emissions and secondary production contribute 28% and 70% of
226 the reduction in the $OVOC^R$, respectively (Supplementary S6).

227 **O_3 mitigation strategy for Chinese cities**

228 Figure 4 shows the combined effect of VOCs and NO_X on normalized O_3 and total
229 OH turnover rate induced by all reactive gases (including VOCs, NO_X and SO_2). The
230 total OH turnover rate is the total reaction rate of all reactive species with OH radical,
231 which closely relates to the level of secondary pollution (Methods M4). The Beijing
232 environment during 2006-2020 was overall in the VOC-limited regime for both O_3 and

233 total OH turnover rate, with a trend of moving toward to the transition regime (between
234 VOC-limited and NO_x-limited regimes). The O₃ trend line parallels the isopleths before
235 2014, followed by efficiently decreasing O₃ after 2014 (Fig. 4A), which is consistent
236 with the observed trend of normalized O₃ as shown in Fig. 1D. Figure 4 clearly shows
237 that the turning point of the O₃ trend appears in 2014 due to a faster decrease in
238 VOC^R/NO_x^R after 2014 than before 2014.

239 If VOCs and NO_x in Beijing maintain their current declining trends over the next
240 five years (blue lines from 2020 to 2025 in Fig. 4), both O₃ and total OH turnover rate
241 will continue to decrease. By 2025 anthropogenic VOCs will have decreased to a very
242 low level and biogenic VOCs will dominate the total VOCs reactivity, which means
243 that further reduction of anthropogenic VOCs may have little impact on VOC^R/NO_x^R.
244 Moreover, in 2025, the O₃ production will have moved to a transition regime. In this
245 case, a faster NO_x reduction after 2025 will help O₃ production rapidly enter a NO_x-
246 limited regime, which favors O₃ mitigation through further reducing NO_x (magenta
247 lines after 2025 in Fig. 4). The situation in Shanghai (Extended data Fig. 6) is similar
248 to that in Beijing. In addition to the two largest cities, most urban regions in China have
249 been changing from a VOC-limited regime to a transition regime^{35,36}. These findings
250 suggest that priority on VOCs abatement in the short term, followed by NO_x abatement
251 in the long term is the optimal strategy for summertime O₃ air quality improvement in
252 Chinese cities.

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257 **Acknowledgements**

258 X. L. received financial support of the National Natural Science Foundation of China
259 (91644108, 91844301). X. L. received financial support of the Beijing Municipal
260 Natural Science Fund for Distinguished Young Scholars (grant No. JQ21030). F. R. and
261 A. W. received financial support of the Federal Ministry of Education and Research,
262 Germany (01LP1929A, Practice). H. S. received financial support of the National Key
263 Scientific and Technological Infrastructure project “Earth System Numerical
264 Simulation Facility” (EarthLab).

265 **Author contributions statement**

266 H. S., X. L., Y. Z., and Y. C. conceived, and designed this study, and revised the article
267 critically. W. W. and X. L. acquired, analyzed, and interpreted data, drafted the article,
268 and revised it critically. R. N. and Y. C. performed the chemical transport model
269 simulations. D. D. P., F. R., A. W., R. N., Z. T. and U. P. revised the article critically. Q.
270 Z. provided the emission inventory data. Y. L., S. L., Y. W., S. C., K. L., M. H., L. Z.,
271 M. S., C. H., X. T., K. M. L., L. C., M. F. collected data.

272 **Competing interest statement**

273 Authors declare no competing interests.

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285 **Figure Legends**

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287 **Figure 1. Variations in ozone and OH reactivity of precursors derived from**
288 **observations in Beijing, August between 2006 and 2020.** Variations in daytime (6:00-
289 19:00) averages of VOC^{R} (including primary hydrocarbons plus carbon monoxide),
290 NO_x^{R} and column concentrations of NO_2 (A), MDA8 O_3 concentrations (B), the ratio
291 of VOC^{R} to NO_x^{R} (C), and normalized MDA8 O_3 concentrations (D). Dashed lines
292 show linear regression fits to observations in periods before and after 2014. Model-
293 simulated MDA8 O_3 concentrations are included in (B).

294

295 **Figure 2. Variations in model-simulated OH concentrations and its driving forces**
296 **in Beijing.** (A) Variations in simulated daytime average OH concentrations (red dots)
297 and measured and simulated (MVK+MACR)/isoprene ratios (blue dots and circles) in
298 Beijing in August between 2006 and 2020. Dashed lines show exponential regression
299 fits to trends. (B) Apportionment of the causes of the change in OH derived from box
300 model sensitivity analysis. (C) Variations in OH turnover rate induced by VOCs
301 ($\text{OH} \times \text{VOC}^{\text{R}}$). Dashed lines show linear regression fits to trends in periods before and
302 after 2014. (D) Correlation of $\text{OH} \times \text{VOC}^{\text{R}}$ with observed MDA8 O_3 and daytime O_x
303 concentrations.

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305 **Figure 3. The impact of the $\text{VOC}^{\text{R}}/\text{NO}_x^{\text{R}}$ ratio on ozone production.** (A) Variations

306 in P(RO_x) and ChL. The three metrics are colored by the VOC^R/NO_x^R ratio. (B) Hourly
307 mean diurnal profiles of modeled rates of primary RO_x production during 2006-2020.
308 (C) Modeled (circle dots) and observed (black diamonds) dependence of OVOC^R on
309 VOC^R/NO_x^R ratio. The modeled results are colored by different VOC^R levels. The
310 observed results indicate mean values of OVOC^R with standard deviations at each
311 VOC^R/NO_x^R range, with sample size of 132 daily average values in total. (D)
312 Mechanism diagram showing the impact of decreased VOC^R/NO_x^R ratio on ozone
313 production.

314

315 **Figure 4. Isopleth plots for (A) ozone and (B) total OH turnover rate as a function**
316 **of NO_x^R and VOC^R.** Red solid circles represent the average levels of MDA8 O₃ and
317 daytime OH turnover rate for each year during 2006-2020, and the solid lines indicate
318 fits to the trends during 2006-2014 and 2014-2020. The yellow dotted lines indicate the
319 VOC^R level from biogenic emissions. The blue and magenta arrows indicate future
320 routes for VOCs and NO_x abatement that mitigate ozone effectively in the short term
321 (2020-2025) and long term (after 2025), respectively.

322 **References:**

- 323 1. Huang RJ, Zhang YL, Bozzetti C, Ho KF, Cao JJ, Han YM, *et al.* High secondary
324 aerosol contribution to particulate pollution during haze events in China. *Nature*
325 2014, **514**(7521): 218-222.
- 326 2. Lu X, Hong JY, Zhang L, Cooper OR, Schultz MG, Xu XB, *et al.* Severe Surface
327 Ozone Pollution in China: A Global Perspective. *Environ Sci Technol Lett* 2018,
328 **5**(8): 487-494.
- 329 3. World Health Organization (WHO), WHO global air quality guidelines: particulate
330 matter (PM_{2.5} and PM₁₀), ozone, nitrogen dioxide, sulfur dioxide and carbon
331 monoxide: executive summary. Publication ISBN: 978-92-4-003443-3, WHO, 2021
332 <https://apps.who.int/iris/bitstream/handle/10665/345334/9789240034433-eng.pdf>.
- 333 4. Zhang Q, Zheng Y, Tong D, Shao M, Wang S, Zhang Y, *et al.* Drivers of improved
334 PM_{2.5} air quality in China from 2013 to 2017. *Proc National Acad Sci* 2019,
335 **116**(49): 24463-24469.
- 336 5. Wang S, Su H, Chen C, Tao W, Streets DG, Lu Z, *et al.* Natural gas shortages during
337 the “coal-to-gas” transition in China have caused a large redistribution of air
338 pollution in winter 2017. *Proc National Acad Sci* 2020, **117**(49): 31018-31025.
- 339 6. Lu X, Zhang L, Wang X, Gao M, Li K, Zhang Y, *et al.* Rapid increases in warm-
340 season surface ozone and resulting health impact in China since 2013. *Environ Sci*
341 *Technol Lett* 2020, **7**(4): 240-247.
- 342 7. Fiore AM, Jacob DJ, Logan JA, Yin JH. Long-term trends in ground level ozone

- 343 over the contiguous United States, 1980–1995. *J Geophys Res-Atmos* 1998,
344 **103(D1):** 1471-1480.
- 345 8. Cooper OR, Parrish D, Ziemke J, Balashov N, Cupeiro M, Galbally I, *et al.* Global
346 distribution and trends of tropospheric ozone: An observation-based review Global
347 distribution and trends of tropospheric ozone. *Elem Sci Anth* 2014, **2:** 29.
- 348 9. Simon H, Reff A, Wells B, Xing J, Frank N. Ozone Trends Across the United States
349 over a Period of Decreasing NO_x and VOC Emissions. *Environ Sci Technol* 2015,
350 **49(1):** 186-195.
- 351 10. Fleming ZL, Doherty RM, Von Schneidemesser E, Malley CS, Cooper OR, Pinto
352 JP, *et al.* Tropospheric Ozone Assessment Report: Present-day ozone distribution
353 and trends relevant to human health. *Elem Sci Anth* 2018, **6:** 12.
- 354 11. Wang W, Parrish DD, Li X, Shao M, Liu Y, Mo Z, *et al.* Exploring the drivers of the
355 increased ozone production in Beijing in summertime during 2005–2016. *Atmos*
356 *Chem Phys* 2020, **20(24):** 15617-15633.
- 357 12. Zhang Q, Yuan B, Shao M, Wang X, Lu S, Lu K, *et al.* Variations of ground-level
358 O₃ and its precursors in Beijing in summertime between 2005 and 2011. *Atmos*
359 *Chem Phys* 2014, **14(12):** 6089-6101.
- 360 13. Liu Y, Wang T. Worsening urban ozone pollution in China from 2013 to 2017 –
361 Part 2: The effects of emission changes and implications for multi-pollutant control.
362 *Atmos Chem Phys* 2020, **20(11):** 6323-6337.
- 363 14. Cheng N, Li R, Xu C, Chen Z, Chen D, Meng F, *et al.* Ground ozone variations at

- 364 an urban and a rural station in Beijing from 2006 to 2017: Trend, meteorological
365 influences and formation regimes. *Journal of Cleaner Production* 2019, **235**: 11-20.
- 366 15. Ma ZQ, Xu J, Quan WJ, Zhang ZY, Lin WL, Xu XB. Significant increase of surface
367 ozone at a rural site, north of eastern China. *Atmos Chem Phys* 2016, **16**(6): 3969-
368 3977.
- 369 16. Monks PS, Archibald A, Colette A, Cooper O, Coyle M, Derwent R, *et al.*
370 Tropospheric ozone and its precursors from the urban to the global scale from air
371 quality to short-lived climate forcer. *Atmos Chem Phys* 2015, **15**(15): 8889-8973.
- 372 17. UN Environment. A review of 20 years' air pollution control in Beijing. Publication
373 ISBN: 978-92-807-3743-1, United Nations Environment Programme, 2019,
374 [https://wedocs.unep.org/bitstream/handle/20.500.11822/27645/airPolCh_EN.pdf?s](https://wedocs.unep.org/bitstream/handle/20.500.11822/27645/airPolCh_EN.pdf?sequence=1&isAllowed=y)
375 [equence=1&isAllowed=y](https://wedocs.unep.org/bitstream/handle/20.500.11822/27645/airPolCh_EN.pdf?sequence=1&isAllowed=y).
- 376 18. Yang Y, Liao H, Lou S. Increase in winter haze over eastern China in recent decades:
377 Roles of variations in meteorological parameters and anthropogenic emissions. *J*
378 *Geophys Res-Atmos* 2016, **121**(21): 13,050-013,065.
- 379 19. Li K, Jacob DJ, Liao H, Shen L, Zhang Q, Bates KH. Anthropogenic drivers of
380 2013–2017 trends in summer surface ozone in China. *Proc National Acad Sci* 2019,
381 **116**(2): 422-427.
- 382 20. Chinese State Council, Three-Year Action Plan on Defending the Blue Sky 2018,
383 http://english.mee.gov.cn/News_service/news_release/201807/t20180713_4466
384 [24.shtml](http://english.mee.gov.cn/News_service/news_release/201807/t20180713_4466).

- 385 21. Kirchner F, Jeanneret F, Clappier A, Krüger B, van den Bergh H, Calpini B. Total
386 VOC reactivity in the planetary boundary layer: 2. A new indicator for determining
387 the sensitivity of the ozone production to VOC and NO_x. *J Geophys Res-Atmos*
388 2001, **106**(D3): 3095-3110.
- 389 22. Sillman S, Logan JA, Wofsy SC. The sensitivity of ozone to nitrogen oxides and
390 hydrocarbons in regional ozone episodes. *J Geophys Res-Atmos* 1990, **95**(D2):
391 1837-1851.
- 392 23. Edwards PM, Brown SS, Roberts JM, Ahmadov R, Banta RM, deGouw JA, *et al.*
393 High winter ozone pollution from carbonyl photolysis in an oil and gas basin.
394 *Nature* 2014, **514**(7522): 351-354.
- 395 24. Sheehy PM, Volkamer R, Molina LT, Molina MJ. Oxidative capacity of the Mexico
396 City atmosphere - Part 2: A RO_x radical cycling perspective. *Atmos Chem Phys*
397 2010, **10**(14): 6993-7008.
- 398 25. Jenkin M, Saunders S, Wagner V, Pilling M. Protocol for the development of the
399 Master Chemical Mechanism, MCM v3 (Part B): tropospheric degradation of
400 aromatic volatile organic compounds. *Atmos Chem Phys* 2003, **3**(1): 181-193.
- 401 26. Womack CC, McDuffie EE, Edwards PM, Bares R, de Gouw JA, Docherty KS, *et*
402 *al.* An Odd Oxygen Framework for Wintertime Ammonium Nitrate Aerosol
403 Pollution in Urban Areas: NO_x and VOC Control as Mitigation Strategies. *Geophys*
404 *Res Lett* 2019, **46**(9): 4971-4979.
- 405 27. Liu YJ, Seco R, Kim S, Guenther AB, Goldstein AH, Keutsch FN, *et al.* Isoprene
406 photo-oxidation products quantify the effect of pollution on hydroxyl radicals over

- 407 Amazonia. *Sci Adv* 2018, **4**(4): 8.
- 408 28. Barkot DJ, Grossenbacher JW, Hurst JM, Shepson PB, Olszyna K, Thornberry T, *et al.*
409 *al.* A study of the NO_x dependence of isoprene oxidation. *J Geophys Res-Atmos*
410 2004, **109**(D11): 12.
- 411 29. Wang W, Li X, Shao M, Hu M, Zeng L, Wu Y, *et al.* The impact of aerosols on
412 photolysis frequencies and ozone production in Beijing during the 4-year period
413 2012–2015. *Atmos Chem Phys* 2019, **19**(14): 9413-9429.
- 414 30. Jacob DJ. Heterogeneous chemistry and tropospheric ozone. *Atmos Environ* 2000,
415 **34**(12-14): 2131-2159.
- 416 31. Kleinman LI, Daum PH, Lee JH, Lee YN, Nunnermacker LJ, Springston SR, *et al.*
417 Dependence of ozone production on NO and hydrocarbons in the troposphere.
418 *Geophys Res Lett* 1997, **24**(18): 2299-2302.
- 419 32. Tonnesen GS, Dennis RL. Analysis of radical propagation efficiency to assess ozone
420 sensitivity to hydrocarbons and NO_x: 1. Local indicators of instantaneous odd
421 oxygen production sensitivity. *J Geophys Res-Atmos* 2000, **105**(D7): 9213-9225.
- 422 33. Volkamer R, Sheehy P, Molina LT, Molina MJ. Oxidative capacity of the Mexico
423 City atmosphere - Part 1: A radical source perspective. *Atmos Chem Phys* 2010,
424 **10**(14): 6969-6991.
- 425 34. Mao J, Ren X, Chen S, Brune WH, Chen Z, Martinez M, *et al.* Atmospheric
426 oxidation capacity in the summer of Houston 2006: Comparison with summer
427 measurements in other metropolitan studies. *Atmos Environ* 2010, **44**(33): 4107-

428 4115.

429 35. Wang W, van der A R, Ding J, van Weele M, Cheng T. Spatial and temporal changes
430 of the ozone sensitivity in China based on satellite and ground-based observations.
431 *Atmos Chem Phys* 2021, **21**(9): 7253-7269.

432 36. Kang M, Zhang J, Zhang H, Ying Q. On the Relevancy of Observed Ozone Increase
433 during COVID-19 Lockdown to Summertime Ozone and PM2.5 Control Policies
434 in China. *Environ Sci Technol Lett* 2021, **8**(4): 289-294.

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454 **Methods**

455 **M1 Measurements**

456 Field measurements were performed at an urban site in Beijing, August between
457 2006 and 2020. The site (39.99° N, 116.31°E) is located on the roof of a six-story
458 building (~ 20 m above the ground level) on the campus of Peking University
459 (PKUERS). Ozone (O₃) nitric oxide (NO) nitrogen dioxide (NO₂), sulfur dioxide (SO₂),
460 carbon monoxide (CO), VOCs, aerosol optical properties, aerosol surface area
461 concentration (S_a), photolysis frequencies of O₃, NO₂, HONO and formaldehyde
462 (HCHO) and meteorological factors including temperature, relative humidity, wind
463 speed, wind direction and air pressure were measured during the study period. VOCs
464 include nonmethane hydrocarbons (NMHCs) with 56 species and C2-C6 OVOCs. For
465 all these gas pollutant measurements, the calibration of instruments was regularly
466 conducted. The measurement techniques are summarized in the Extended data Table 1.
467 In addition, OMI tropospheric NO₂ columns and HCHO columns were also used for
468 analysis (see details in Supplementary S1). Planetary boundary layer height in Beijing
469 was obtained from the NOAA Air Resource Laboratory website
470 (<https://ready.arl.noaa.gov/READYamet.php>).

471 Figure S1 shows a comparison between the PKUERS site and the other 12 sites
472 from the national monitoring network (NMN) in Beijing (<https://quotsoft.net/air/>, NMN
473 data available from 2014). The O₃ concentration exhibits very similar trends at the
474 PKUERS site and the NMN sites. We have also looked at the O₃ precursor NO₂ (VOC

475 measurements are not available at the NMN sites), which exhibits also similar
476 decreasing trends at both the NMN sites and the PKUERS site. These results show that
477 data from the PKUERS site can represent the long-term trends of both O₃ and its
478 precursor NO₂ in Beijing.

479 In addition to data set in Beijing, data collected at sites in Shanghai and Hong
480 Kong were also used for additional support. Field measurements of trace gases (O₃,
481 VOCs, NO_x, CO, SO₂), aerosols, photolysis frequencies and meteorological
482 parameters were performed at typical urban sites in Shanghai between 2011 and 2019
483 and Hongkong between 2013 and 2019. Relevant measurement data were analyzed for
484 comparison with Beijing. For all analysis, we eliminated the days with abnormally high
485 cloud optical depth when O₃ concentrations were extremely low.

486 **M2 Multiple linear regression**

487 We applied a multiple linear regression (MLR) method to quantify the effect of
488 meteorological and other physical factors on O₃:

$$489 \quad [O_3] = \beta_0 + \sum_{i=1}^n \beta_i X_i + \text{interaction terms} \quad (1)$$

$$490 \quad \text{interaction terms} = \sum \beta_{i,j} X_i X_j, \text{ for } 1 \leq i, j \leq n, \quad (2)$$

491 where [O₃] is the daily MDA8 O₃ concentration and (X₁, ..., X_n) are major
492 meteorological and physical variables. The interaction terms are up to second order,
493 including both X_i² and X_iX_j (i≠j). Similar MLR methods have been successfully applied to
494 quantify the effect of meteorological variability on air pollutants in North America,
495 Europe, and China^{19, 37, 38}. Seven variables (j(O¹D), temperature, relative humidity,
496 wind speed, wind direction, air pressure and boundary layer height) are candidates for

497 the regression. According to the fitting test, $j(O^1D)$, temperature, relative humidity,
498 wind speed and wind direction are more important than air pressure and boundary layer
499 height in determining the coefficient of determination (R^2). Accordingly, the five
500 dominant meteorological variables were selected for the regression. The lag terms (i.e.,
501 the effect of July conditions on August ozone) have a small effect on R^2 , thus they were
502 not considered in this study.

503 Hourly values of O_3 were used to calculate daily MDA8 O_3 concentrations, and
504 hourly values of meteorological factors were averaged to daily averages. Daily averages
505 of the five dominant meteorological variables were regressed onto daily MDA8 O_3
506 concentrations to fit the effect of meteorological variability on O_3 . The coefficient of
507 determination (R^2) for the MLR model is 0.65. Then the fitting results were averaged
508 to acquire the monthly O_3 anomaly of each year induced by meteorological variability.
509 The effect of meteorological factors on O_3 can be normalized by subtracting the
510 monthly O_3 anomaly from observed monthly O_3 concentrations.

511 Notably, the decreases of $PM_{2.5}$ over this 15-year period resulted in increases of
512 actinic flux and photolysis frequencies (**Fig. S2 and Fig. S5**), which could increase O_3
513 production^{11,29}. According to the model simulation of Wang et al¹¹, the rapid increase
514 in actinic flux causes an increase in MDA8 O_3 by 1.6 ppbv (2.3%) yr^{-1} during 2006-
515 2014, accounting for 66% of the increase in O_3 concentrations during this period. In
516 this study, we aim to explore emission control strategy for O_3 mitigation by
517 investigating the relationship between O_3 trend and precursor trend. In this case, we
518 must remove the effect of the change in photolysis frequencies on the O_3 trend by the

519 normalization of $j(\text{O}^1\text{D})$, which helps to show a clearer relationship between O_3 trend
520 and precursor trend.

521 **M3 Model simulation**

522 The photochemistry of radicals and ozone in Beijing were studied by using a zero-
523 dimension photochemical box model. Previous studies indicate that the variation in
524 ozone concentrations at PKUERS site was mainly influenced by local photochemistry,
525 and background ozone contributed < 30% of the variation in ozone concentrations ^{11,12}.
526 In this case, box model is suitable for studying the photochemistry of ozone and OH
527 radicals at PKUERS site. The box model has the advantage of allowing detailed
528 treatment of the VOC oxidation chemistry, at the expense of a comprehensive
529 representation of dynamical processes.

530 The photochemical box model uses the Master Chemical Mechanism (MCM)
531 v3.3.1 as chemical mechanism ³⁹. The MCMv3.3.1 chemistry scheme contains 5832
532 species and 17224 reactions, which can accurately simulate the radical source,
533 conversion and termination to understand the O_3 formation. Model calculations are
534 constrained to measured meteorological factors (i.e., photolysis frequencies,
535 temperature, relative humidity and air pressure), inorganic species (NO , NO_2 , CO , O_3
536 and SO_2), NMHCs and aerosol surface area concentrations. The OVOC species that
537 were continuously measured during the 15 years including acetaldehyde, acetone,
538 methyl ethyl ketone, MVK and MACR were constrained in the model. HCHO was not
539 constrained in the model due to the limitation of measurement data and thus was

540 simulated by the model. To consider the contribution of primary emissions to HCHO,
541 we set emission rates of HCHO in Beijing in the model ($3.3\text{-}11.5\times 10^{-12}$ kg m⁻² s⁻¹),
542 which is derived from local emission inventories - the MEIC inventory
543 (<http://www.meicmodel.org/>). To evaluate the model performance, the model-simulated
544 HCHO was compared with observations in August 2008, 2011 and 2012, which shows
545 a difference of 14% for daytime averages (Fig. S10). This difference is smaller than the
546 uncertainty of model simulation (30%) associated with the uncertainty of various
547 measurement parameters, reaction rate constants and emission rates of HCHO
548 (Supplementary S7), suggesting a good performance of the model in the simulation of
549 HCHO. The five constrained OVOCs and HCHO totally contribute >65% of total
550 OVOC^R and >70% of total P(RO_x) from OVOC photolysis (Fig. S11).

551 The model runs were performed in a time-dependent mode with a 1-hour
552 resolution. A whole month (August) of each year was simulated with three days' spin-
553 up. The diurnal variation in boundary layer height was expressed as a dilution term in
554 the box model to represent the physical losses. The convection of O₃ between near
555 surface air and the residual layer aloft was considered in the box model, according to
556 the method of Womack et al.²⁶. The dry deposition rate for O₃ was set as 0.42 cm s⁻¹ in
557 the daytime and 0.14 cm s⁻¹ in the nighttime, as the deposition rate is known to decrease
558 after sunset^{26, 40, 41}. To account for background O₃ transported from outside the
559 investigated cities, we set the background O₃ of 35 ppbv, which are derived from
560 observed O₃ concentrations at regional background sites in China⁴². For other species
561 such as radicals and OVOCs, the concentrations mixed into the box was set as zero

562 because of their short lifetime.

563 Aerosols can influence O₃ production by heterogeneous reactions such as uptake
564 of HO₂, N₂O₅, NO₂ and NO₃. This includes the uptake coefficients are 0.08 for
565 conversion of HO₂ to H₂O⁴³, 0.007 for conversion of N₂O₅ to HNO₃⁴⁴, 2×10^{-5} for
566 conversion of NO₂ to HONO and HNO₃ (which yields a good simulation of
567 HONO/NO₂ concentration ratios in Beijing⁴⁵) and 1×10^{-3} for conversion of NO₃ to
568 HNO₃³⁰.

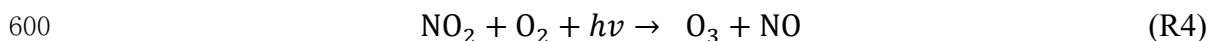
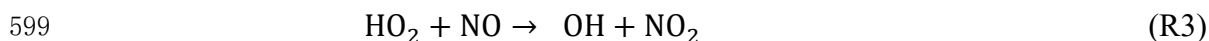
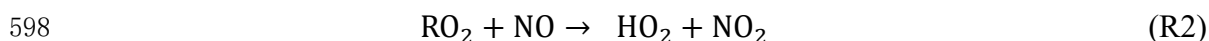
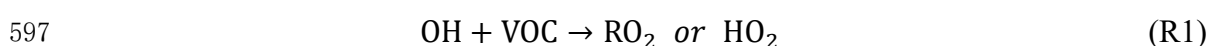
569 Direct measurements of OH concentration were not made during the study period;
570 hence, we simulated OH concentrations by the box model. To identify the primary
571 causes of the continuous OH increase from 2006 to 2020, we performed sensitivity
572 analyses of the relevant factors. The investigated factors include photolysis frequencies,
573 NO_x, VOCs, aerosol uptake and other factors including O₃, H₂O and temperature. We
574 take the simulation result in 2006 as the base scenario. For the testing scenario, we set
575 one of investigated factors to the level of 2020, with other conditions fixed to levels of
576 2006. The difference in simulated OH concentrations between the base scenario and the
577 testing scenario represents the effect of this investigated factor on the OH increase.
578 Notably the effect of aerosol uptake is quantified by changing aerosol surface area
579 concentrations in the model. Similar sensitivity analysis method has been applied by Li
580 et al.¹⁹

581 Due to the limitation of measurement data, some important OVOCs are not
582 constrained in the model such as acrolein, n-butanal, n-hexanal, n-pentanal, propanal,
583 pentanone, methanol and C7-C10 OVOCs. We investigated the impact of whether these

584 OVOCs are constrained in the model or not on our analysis. The model simulation
 585 focuses on August 2017 when these OVOCs were measured. As shown in Fig. S14,
 586 daytime averages of O₃ concentrations, P(O₃), P(RO_x) and ChL changed by 10%, 17%
 587 and 12% and 11% between the two scenarios, indicating that our modelling setup
 588 without these OVOCs constrained is acceptable for the analysis of this study although
 589 some bias exists. Even though, the impact of these unconstrained OVOCs on O₃
 590 production and trends cannot be ignored, which should be paid more attention to in the
 591 future. In addition, other important photodegradable OVOCs that were not measured
 592 such as peroxyacetic acid and pyruvic acid may also cause bias.

593 **M4 Calculation of P(O₃), P(RO_x), ChL and total OH turnover rate**

594 The dominant chemical source of O₃ in the troposphere is the reaction of OH with
 595 VOCs to produce peroxy radicals (HO₂ and RO₂), which then react with NO to yield
 596 NO₂⁴⁶. The overall process can be summarized as:



601 The ozone production rate (P(O₃)) is equal to the reaction rate of peroxy radicals
 602 with NO (i.e., sum of rates of R2 and R3):

$$603 \quad P(\text{O}_3) = k_{\text{HO}_2+\text{NO}}[\text{HO}_2][\text{NO}] + \sum_i(k_{\text{RO}_2+\text{NO}}^i [\text{RO}_2^i][\text{NO}]) \quad (3)$$

604 By defining an average reaction rate constant for all RO₂+NO reactions, $k_{\text{RO}_2+\text{NO}}$,
 605 we obtain:

606
$$P(O_3) = k_{HO_2+NO}[HO_2][NO] + k_{RO_2+NO}[RO_2][NO] \quad (4)$$

607 Our model indicates that k_{RO_2+NO} was in the range of $8.45\text{--}9.58 \times 10^{-12} \text{ cm}^3$
 608 $\text{molecule}^{-1} \text{ s}^{-1}$ in Beijing 2006-2020, which was close to k_{HO_2+NO} value (8.45×10^{-12}
 609 $\text{cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$). Thus:

610
$$P(O_3) \approx k_{HO_2+NO}([HO_2][NO] + [RO_2][NO]) = \alpha k_{HO_2+NO}[HO_2][NO]$$

 611 (5)

612
$$\alpha = ([HO_2] + [RO_2])/[HO_2] \quad (6)$$

613 The model simulation indicates that the $\alpha \approx 1.8$ in Beijing, which remained
 614 approximately constant over the 2006-2020 period. The uncertainty in various
 615 measurement parameters and reaction rate constants leads to 22% uncertainty of α
 616 (Supplementary S7). To test the effect of un-included OVOCs with larger carbon
 617 number on α , we increased C7-C10 OVOCs by 0.68 ppb (the observed value in 2017)
 618 in the model, resulting in an increase of α by 5%.

619 ChL characterizes the number of iterations each RO_X radical makes prior to
 620 termination, and is the ratio of OH cycling rate (i.e., the rate of reaction R3) to radical
 621 primary production rate ($P(RO_X)$). ChL depends on the levels of VOCs and NO_X with
 622 modest impact from actinic flux.

623
$$\text{ChL} = \alpha \frac{k_{HO_2+NO}[HO_2][NO]}{P(RO_X)} \quad (7)$$

624 The primary production of the peroxy radicals, $P(RO_X)$, is determined by the
 625 photolysis of radical precursors (including carbonyls, nitrous acid and O_3) and alkene
 626 ozonolysis; these are the radical chain initiation reactions:

627
$$P(RO_X) = 2 \times [O_3] \times j(O^1D) \times \theta + [HONO] \times j(HONO)$$

 628
$$+ \sum_i ([OVOC_i] \times j_i \times \gamma_i) + \text{alkene ozonolysis} \quad (8)$$

629 where θ is the fraction of O^1D from ozone photolysis that reacts with H_2O . $OVOC_i$

630 represents each OVOCs species, j_i represents the photolysis frequency of each OVOC_i,
 631 and γ_i represents the number of RO_x radicals generated from the photolysis of each
 632 OVOC_i molecule. For most OVOCs species, γ_i is equal to 2 or 0.

633 When the radical chain cycling is much more important than chain initiation (or
 634 equivalently, chain termination), i.e., ChL \gg 1, R1 and R3 dominate the source and sink
 635 of peroxy radicals (HO₂ and RO₂), respectively. In this case, the reaction rates of R1
 636 and R3 are approximately equal given the short lifetime of peroxy radicals (~ several
 637 minutes). **Extended data Fig. 5** shows that $k_{\text{HO}_2+\text{NO}}[\text{HO}_2][\text{NO}]:[\text{OH}]\times\text{VOC}_R$ gets
 638 closer to 1:1 as ChL increases to 5. In Beijing, the ChL was in the range of 4.3~5.4 (**Fig.**
 639 **3A**). Under this condition, the reaction rates of R1 and R3 were approximately equal:

$$\begin{aligned}
 640 \quad P(O_3) &= \alpha k_{\text{HO}_2+\text{NO}}[\text{HO}_2][\text{NO}] \\
 641 \quad &= \alpha k_{\text{OH}+\text{VOC}}[\text{OH}][\text{VOC}] = \alpha[\text{OH}]\text{VOC}_R \quad (9)
 \end{aligned}$$

642 The total OH turnover rate is the total reaction rate of all reactive species with OH
 643 radical:

$$\begin{aligned}
 644 \quad \textit{Total OH turnover rate} &= [\text{OH}]\times\sum_j k_{\text{OH}+S_j} [S_j] \\
 645 \quad (10)
 \end{aligned}$$

646 $\sum_j k_{\text{OH}+S_j} S_j$ is the total OH reactivity of all reactive species including VOCs,
 647 NO_x, SO₂ and CO. S_j represents each reactivity species. $k_{\text{OH}+S_j}$ is the reaction rate
 648 coefficient for the reaction between reactive species S_j and OH. The total OH turnover
 649 rate characterizes the rate of secondary production (including ozone and secondary
 650 aerosols) induced by OH, which closely relates to the level of secondary pollution.

651 O₃ precursor sensitivity depends on the dominant loss pathways of RO_x radicals.
 652 O₃ production is NO_x-limited if the self-reaction of peroxy radicals dominates the RO_x

653 sink, and VOC-limited if the reaction of NO₂ with OH dominates^{31,46}. Accordingly, the
 654 ratio of OH+NO₂ reaction rate to the total rates of the two RO_X sinks, i.e., L_N/Q, is used
 655 to identify O₃ sensitivity regimes. O₃ production is NO_X-limited if L_N/Q is lower than
 656 0.5, otherwise, it is VOC-limited or NO_X-saturated³¹. As shown in Fig. S15, L_N/Q is
 657 larger than 0.5 from 2006 to 2020, indicating a NO_X-saturated regime.

658
$$L_N/Q = \frac{k_{OH+NO_2}[OH][NO_2]}{k_{HO_2+RO_2}[HO_2][RO_2]+k_{HO_2+HO_2}[HO_2][HO_2]+k_{OH+NO_2}[OH][NO_2]} \quad (11)$$

659

660 **Data availability**

661 The data and code to generate the results in the manuscript are freely available at Open
662 Research Data Repository of the Max Planck Society (<https://doi.org/10.17617/3.LEFS4A>).

663 **Code availability**

664 The box model is available from <https://sites.google.com/site/wolfegm/models>.

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667 **Methods-only References**

668 37. Otero N, Sillmann J, Mar KA, Rust HW, Solberg S, Andersson C, *et al.* A multi-model
669 comparison of meteorological drivers of surface ozone over Europe. *Atmos Chem Phys*
670 2018, **18**(16): 12269-12288.

671 38. Tai APK, Mickley LJ, Jacob DJ. Correlations between fine particulate matter (PM_{2.5})
672 and meteorological variables in the United States: Implications for the sensitivity of
673 PM_{2.5} to climate change. *Atmos Environ* 2010, **44**(32): 3976-3984.

674 39. Wolfe GM, Marvin MR, Roberts SJ, Travis KR, Liao J. The framework for 0-D
675 atmospheric modeling (F0AM) v3. 1. *Geosci Model Dev* 2016, **9**(9): 3309.

676 40. Zhang L, Brook JR, Vet R. A revised parameterization for gaseous dry deposition in air-
677 quality models. *Atmos Chem Phys* 2003, **3**: 2067-2082.

678 41. Yang S, Yuan B, Peng Y, Huang S, Chen W, Hu W, *et al.* The formation and mitigation
679 of nitrate pollution: comparison between urban and suburban environments. *Atmos*
680 *Chem Phys* 2022, **22**(7): 4539-4556.

- 681 42. Xu X, Lin W, Xu W, Jin J, Wang Y, Zhang G, *et al.* Long-term changes of regional
682 ozone in China: implications for human health and ecosystem impacts. *Elem Sci Anth*
683 2020, **8**.
- 684 43. Tan Z, Hofzumahaus A, Lu K, Brown SS, Holland F, Huey LG, *et al.* No Evidence for
685 a Significant Impact of Heterogeneous Chemistry on Radical Concentrations in the
686 North China Plain in Summer 2014. *Environ Sci Technol* 2020, **54**(10): 5973-5979.
- 687 44. Wang H, Lu K, Tan Z, Sun K, Li X, Hu M, *et al.* Model simulation of NO₃, N₂O₅ and
688 ClNO₂ at a rural site in Beijing during CAREBeijing-2006. *Atmos Res* 2017, **196**: 97-
689 107.
- 690 45. Hendrick F, Muller JF, Clemer K, Wang P, De Maziere M, Fayt C, *et al.* Four years of
691 ground-based MAX-DOAS observations of HONO and NO₂ in the Beijing area.
692 *Atmospheric Chemistry and Physics* 2014, **14**(2): 765-781.
- 693 46. Kleinman LI. The dependence of tropospheric ozone production rate on ozone
694 precursors. *Atmos Environ* 2005, **39**(3): 575-586.
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